CHAPTER 4

STRUCTURAL RESTORATION

2 Martin Stark

Seismic Reconstruction, Thermal and Maturity Modeling of the Nova Scotia – Morocco Conjugate Margins - 2019



PL. 4.1.A

Seismic Reconstruction, Thermal and Maturity Modeling of the Nova Scotia – Morocco Conjugate Margins - 2019



PL. 4.1.B

Sequential restoration of the T1 conjugate transect

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PL. 4.2.A

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PL. 4.2.B

Sequential restoration of the T2 conjugate transect

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Seismic Reconstruction, Thermal and Maturity Modeling of the Nova Scotia – Morocco Conjugate Margins - 2019 NORTHERN MOROCCO NW SE ~2137 km 90,000 100,000 50,000 20.000 50 OOF 60.000 70 000 80.000 60,000 40,000 30,000 ~1470 km 60.00 50,000 40,000 30.000 80,000 90,00 ~1080 km ~320 km --2,000 --3,000 --4,000 --5,000 --6,000 --7,000 --7,000 --8,000 --9,000 --10,000



PL. 4.3.B

Sequential restoration of the T3 conjugate transect

200Ma

SE

101Ma

130Ma

145Ma

163Ma

183Ma

Upper Jurassic carbonate platforms Mostly vertical diapirism stopping end of Cretaceous (probably different causes)

BOTH MARGINS

Vertical exaggeration x2

9,000

50 km

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PL. 4.4.A

Seismic Reconstruction, Thermal and Maturity Modeling of the Nova Scotia – Morocco Conjugate Margins - 2019



Sequential restoration of the T4 conjugate transect

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Nova Scotia margin

Following the sub-division proposed by Shimeld, 2004; Albertz et al., 2010; and PFA, 2011 the four cross-sections restored in the present study are located in various sub-provinces and show examples of the different structural styles and salt geometries characterizing the slope domain of the Nova Scotia margin.

Transect 1 is located in the Banquerau Syn-kinematic Wedge province described in detail by Ings and Shimeld, 2006. It is characterized by a major gravity gliding allochthonous wedge active mainly during the Jurassic to Early Cretaceous. This wedge is detached on top of a large early-created salt tongue. Late vertical diapirs define the southwest boundaries of this major wedge.

Transect 2 goes through the Canopy province described in detail by Deptuck et al; 2009; Deptuck, 2010; Deptuck et al., 2015. It is characterized by large allochthonous salt tongues (canopies) fed by one or several deep salt stems that may be of allochthonous or autochthonous origin. These canopies can be completely disconnected from their feeding zone, the stems corresponding to steeply dipping weld zones.

Transect 3 lies between the Canopy and Diapir provinces (PFA, 2011). Transect 4 is located in the Diapir province characterized by mainly vertically raised diapirs on top of the autochthonous salt basin with local salt tongues of relatively reduced extent (compared to structures in the Canopy province) at the southwestern limit of the salt basin (Deptuck et al. 2010).

General observations

There is a significant variation of sediment thicknesses from the NE to the SW (i.e. from the T1 to the T4). This variation is not the same in the two margins: overall sediment thickness increases from the NE to the SW on the Morocco side whereas it decreases on the Nova Scotia side. This difference should be related to different geological causes, for example the strong sediment supply from the Laurentian River at the north of the Nova Scotia margin, and the presence of the Canary Island Arc capturing sediments starting in the Early Paleogene in the south of the Morocco margin.

Different structural style and salt geometries are noted in the two margins (for example the presence of the Banquereau Syn-kinematic Wedge and the "roho" system in the Canadian margin), which are likely to be mainly related to the initial salt thickness and the sediment supply through time.

In both margins, the salt basin termination to the south has been sampled at level of Transect 4 as described in literature. Both margins are characterized in their conjugate southern parts (Transects 3 and 4) by the presence of Upper Jurassic carbonate platforms.

Relative distribution of the gas and oil provinces related to burial over time should not be comparable between the two margins and even reversed in a N-S direction, as the sediment thicknesses show significant variations.

The Early Jurassic series deposition (including a possible Early Jurassic source rock) is related to mini-basin development controlled by early salt movements. Therefore mini-basin evolution can locally control the distribution and characteristics of the Liassic source rocks.



Based on the Morocco salt basin sub-division proposed by Tari et al., 2003; Tari and Molnar, 2005; Tari and Jabour, 2013, the four crosssections restored in the present study are located in various sub-provinces and show examples of the different structural styles and salt geometries characterizing the slope domain of the northern Morocco margin.

Transect 1 is located in the Safi Sub-basin (Essaouira Basin). It shows the characteristic halokinetic elements of a passive margin such as extensional salt structures beneath the shelf and the upper slope, and compressional features downdip on the lower slope (Tari and Jabour, 2013).

Transect 2 is located in the Agadir Basin characterized by simple, mostly vertical diapirism driven dominantly by the load of supra-salt, mostly Cenozoic age sediments (Tari et al. 2003). Gravity-driven compressional features, such as toe-thrust, are present at the western boundary of the salt basin.

Transect 3 and Transect 4 are located in the Tarfaya Basin and Cap Juby area, respectively, where there are only a few salt pillows and diapiric salt walls outboard of the prominent Jurassic shelf margin (Hafid et al., 2008). The Cap Juby area is bounded to the west by Canary Island volcanism starting in the Paleocene-Eocene.

Implications for petroleum systems

Clastic reservoir presence and distribution can be related to sediment supply and therefore are reversed in the Nova Scotia and northern Morocco margins.

On the other hand, potential Jurassic carbonate reservoirs appear to be similar in the southern part of both margins (T3-T4 areas).